

Multidecadal Tropical Cloud Variability in Observations and GCMs

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Observations

- upper-level (mid+high) cloud cover obtained from surface synoptic cloud reports in EECRA/COADS
- wind components from COADS corrected for linear increase in tropical mean scalar wind speed
- wind components estimated from Smith and Reynolds reconstructed SLP
- Hulme precipitation (continents and islands only)
- Kaplan SST

all data are averaged to seasonal $5^{\circ} \times 10^{\circ}$ means

NCAR CCSM3 Simulations

- 10-member AMIP ensemble (historical SST boundary condition with and without historical radiative forcing)
- 5-member IPCC “20th century” ensemble (coupled model with historical radiative forcing)
- 520-year control run (coupled model with 1990 radiative forcing)

high-level cloud amount and lowest layer wind components are converted to seasonal $5^{\circ} \times 10^{\circ}$ means

GFDL CM2.0 and CM2.1 Simulations

- 3-member IPCC “20th century” ensemble for CM2.0 and for CM2.1
- 480-year control run (1860 radiative forcing) for CM2.0 and for CM2.1

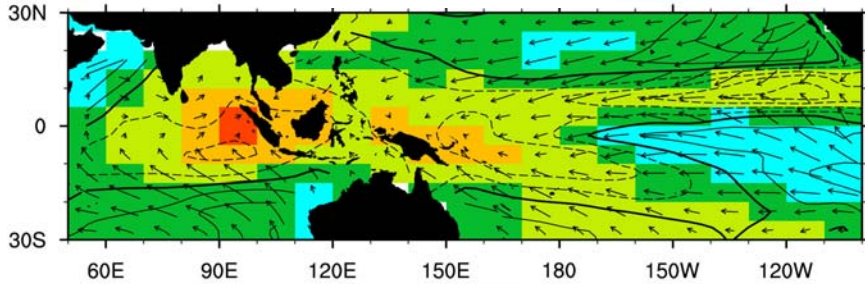
high-level cloud amount and lowest layer wind components are converted to seasonal $5^{\circ} \times 10^{\circ}$ means

CM2.0 and CM2.1 Differences

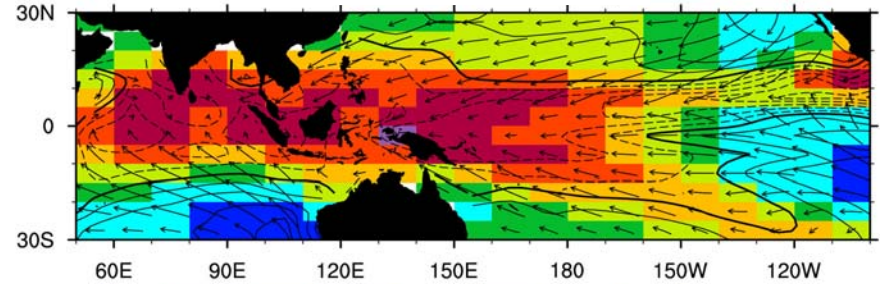
- Atmospheric dynamical core is B-grid for CM2.0 and finite volume for CM2.1
- CM2.1 has a 25% smaller cloud droplet radius threshold for raindrop formation
- CM2.1 has an 18% faster rate of cloud erosion in convective conditions

Cloud Climatologies

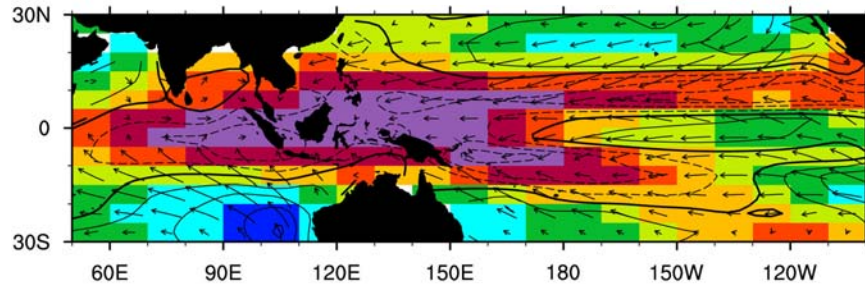
Observations



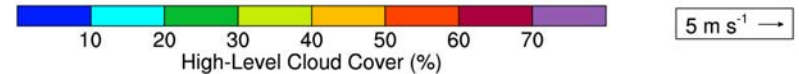
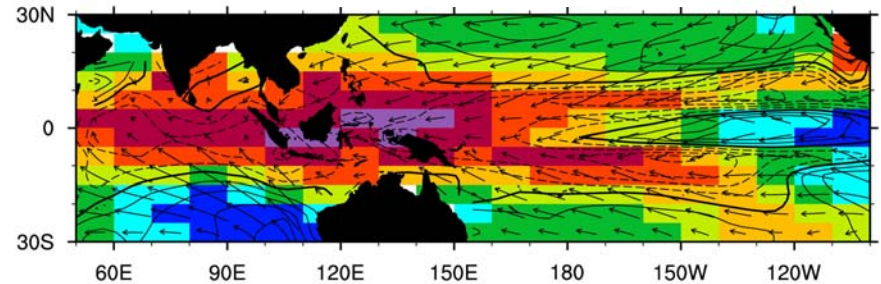
CCSM3 AMIP



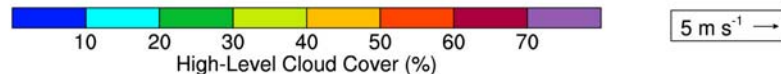
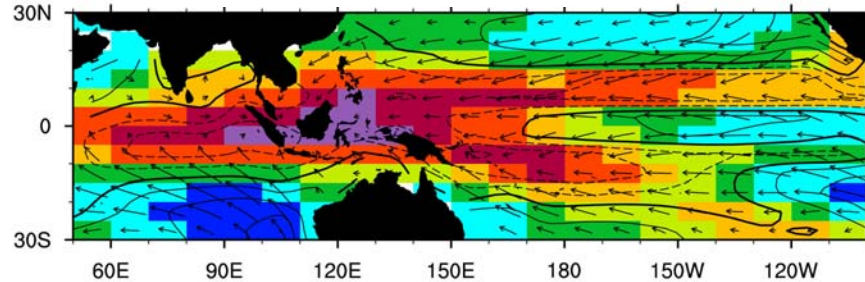
CM2.0 20th Century



CCSM3 20th Century



CM2.1 20th Century



color = cloud cover
arrows = near-surface wind
lines = divergence (every $1 \times 10^{-6} \text{ s}^{-1}$)

Analysis Method

- Relative changes in cloud cover are compared since climatological cloud cover substantially differs between observations and simulations.
- The 1957-1976 average is subtracted from the 1977-1996 average for observations, AMIP, and IPCC 20th century simulations.

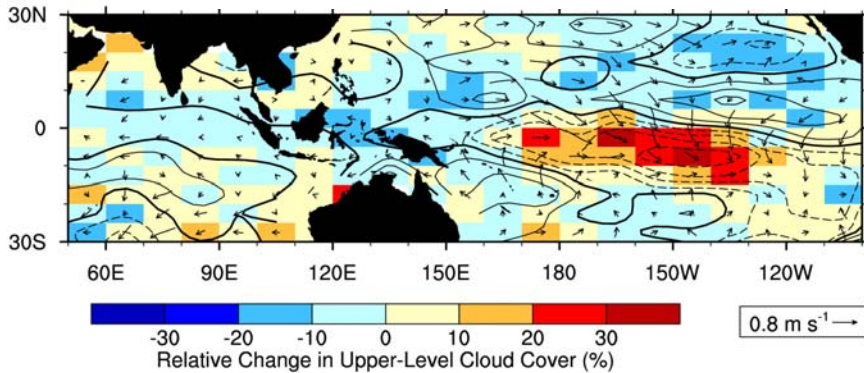
Do the GCMs reproduce the observed pattern and magnitude of change?

- Running differences between 20-year running averages are calculated for the control runs.

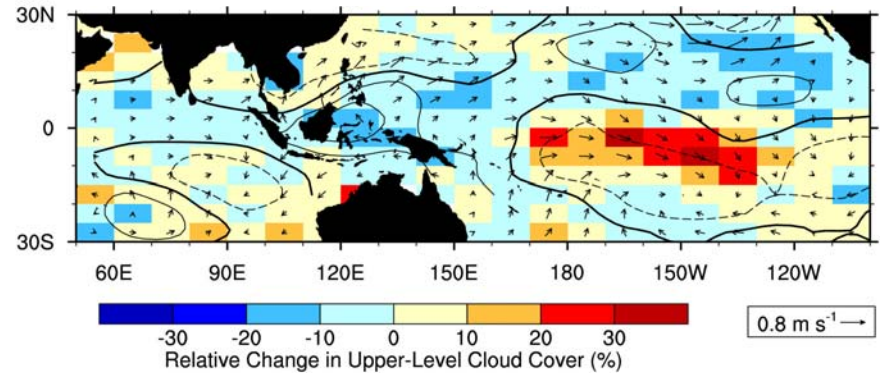
Is natural variability in the GCMs as large as the observed change?

Observed Interdecadal Change

Cloud and COADS Divergence

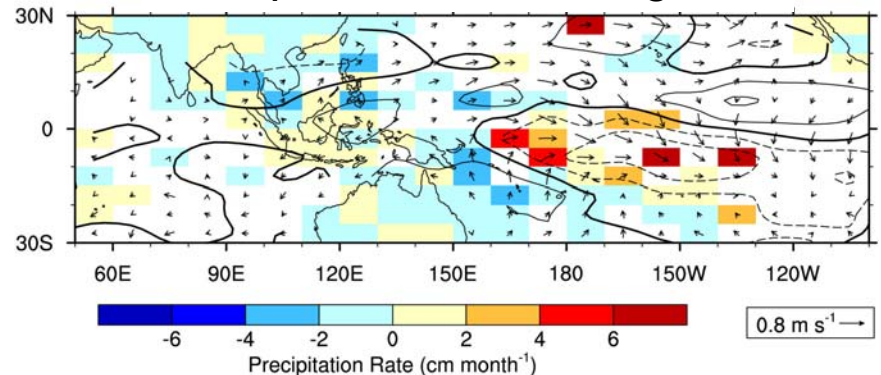


Cloud and SLP-derived Divergence



color = cloud cover or precipitation rate
arrows = near-surface wind
lines = divergence (every $0.15 \times 10^{-6} \text{ s}^{-1}$)

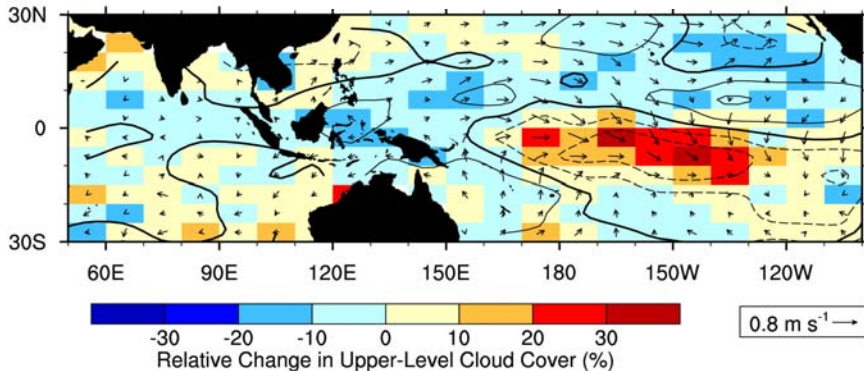
Precipitation and Divergence



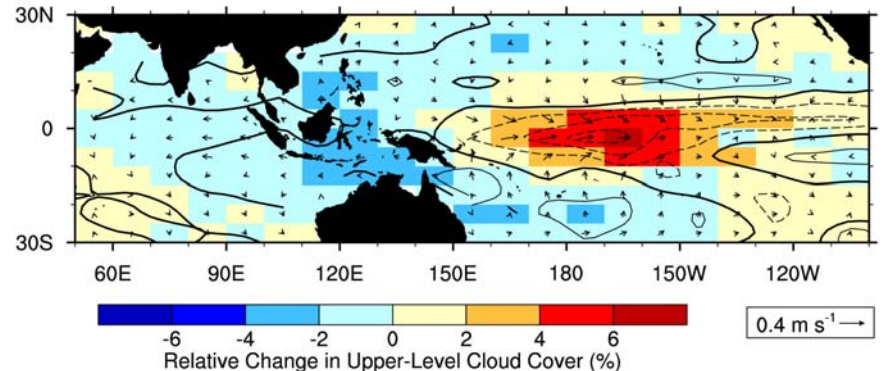
upper-level cloud cover and precipitation have generally increased where there is more convergence and decreased where there is more divergence

Not Merely Due to Interannual El Niño

Actual Change



Change Due to Regression on Niño3.4



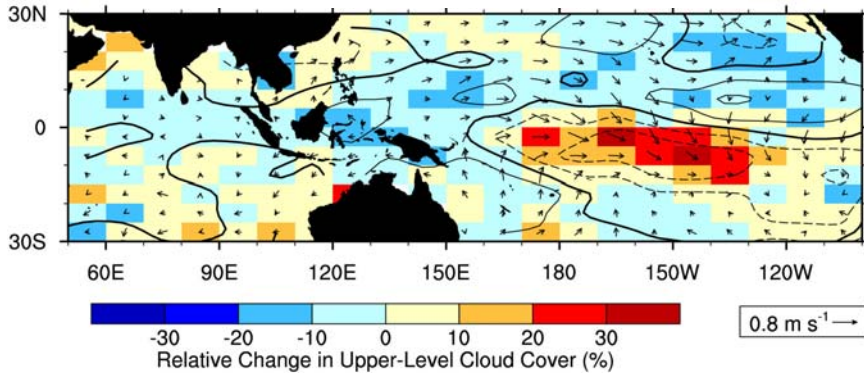
color = cloud cover (every 10%)
arrows = near-surface wind (0.8 m s^{-1})
lines = divergence (every $0.15 \times 10^{-6} \text{ s}^{-1}$)

color = cloud cover (every 2%)
arrows = near-surface wind (0.4 m s^{-1})
lines = divergence (every $0.08 \times 10^{-6} \text{ s}^{-1}$)

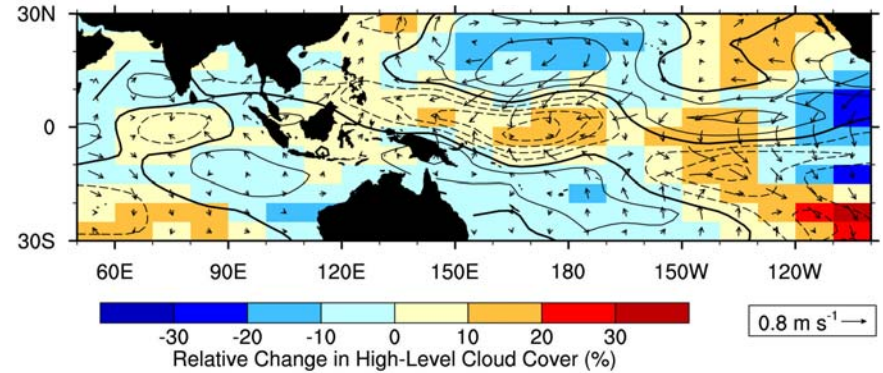
the cloud cover change accounted for by a linear relationship to interannual El Niño variations has a similar pattern but much smaller magnitude than the interdecadal change

AMIP Comparison

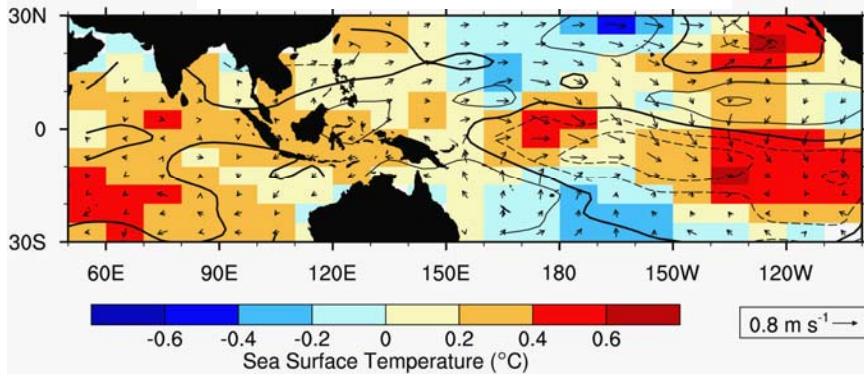
Observed Cloud Change



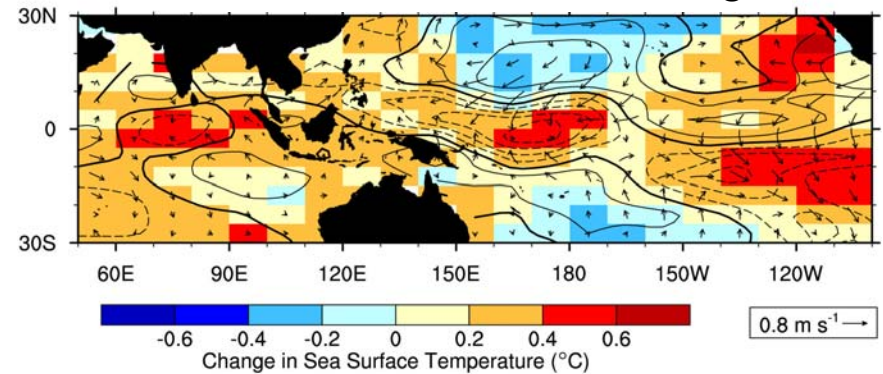
CCSM3 AMIP Cloud Change



Observed SST Change



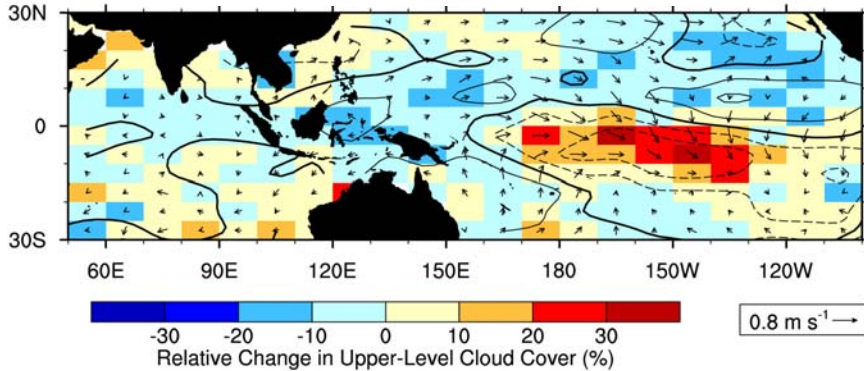
CCSM3 AMIP SST Change



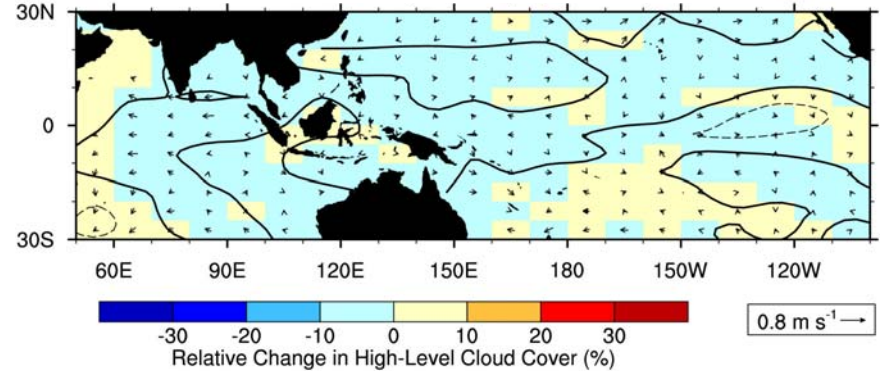
upper-level cloud cover changes in AMIP runs are weaker than observed changes and have a generally similar pattern, albeit more localized on SST anomalies

20th Century Run Comparison

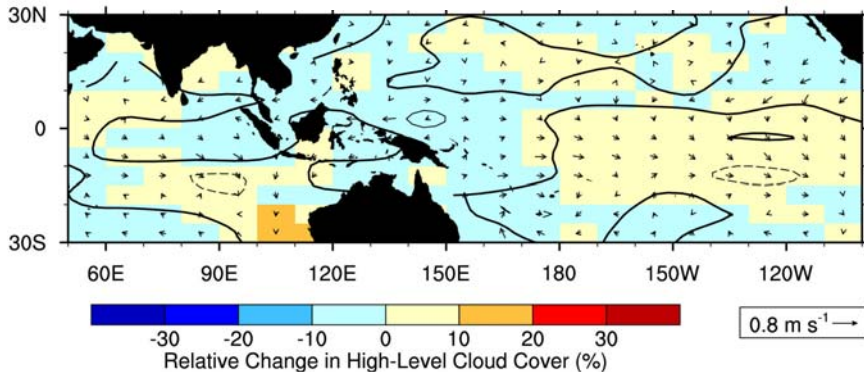
Observed Cloud Change



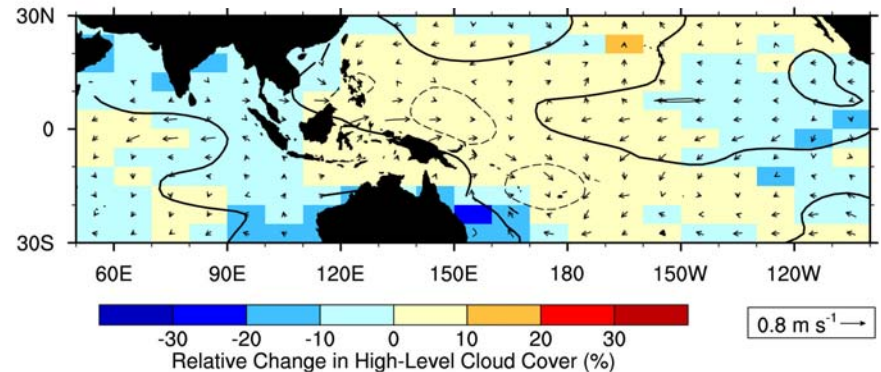
CCSM3 Cloud Change



CM2.0 Cloud Change



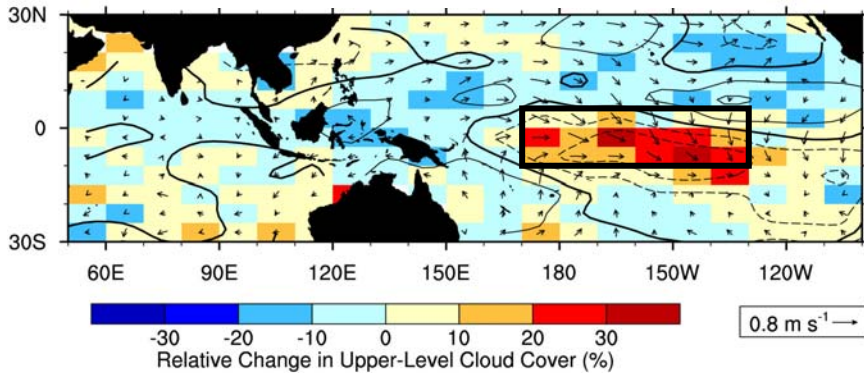
CM2.1 Cloud Change



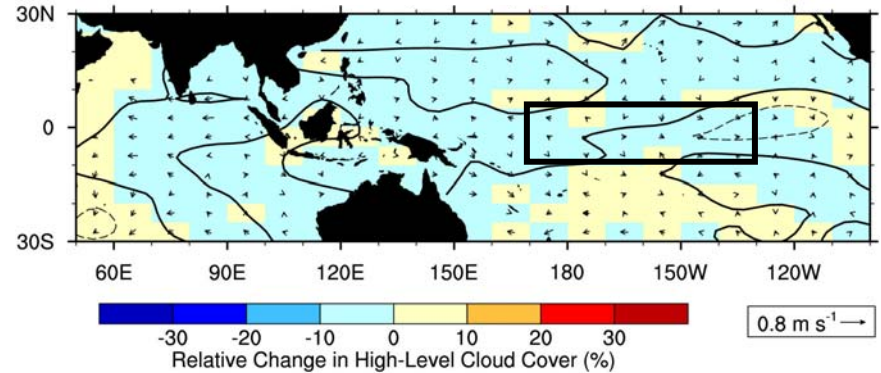
upper-level cloud cover changes in the 20th Century runs do not resemble the observed changes because the evolution of the coupled atmosphere-ocean model is different from what happened in the real world

Averaging Region for Control Comparison

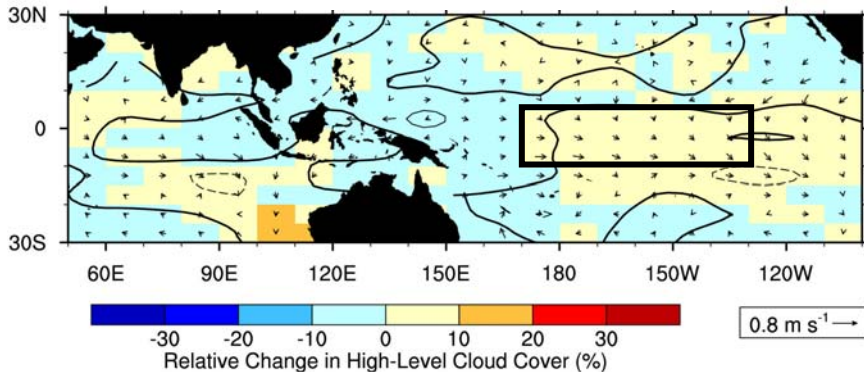
Observed Cloud Change



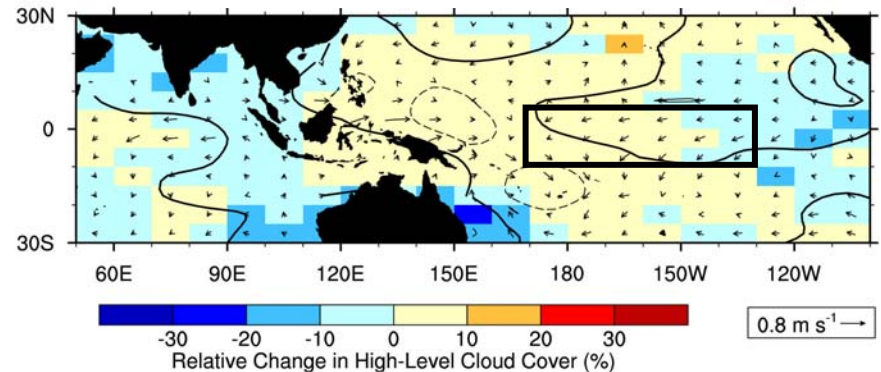
CCSM3 Cloud Change



CM2.0 Cloud Change



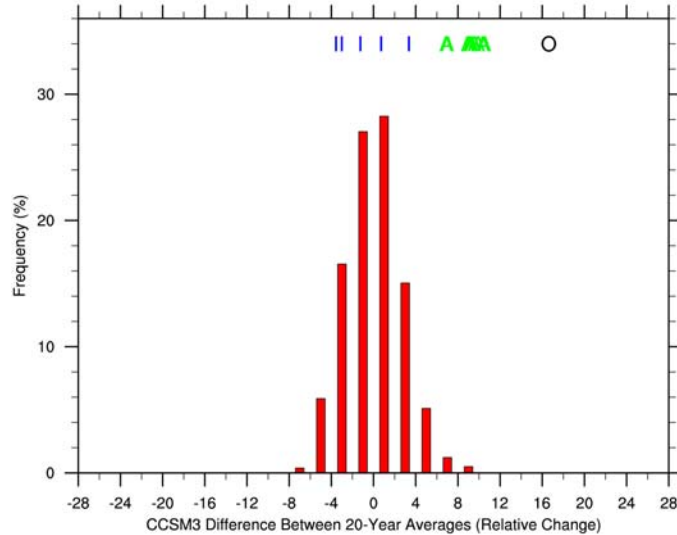
CM2.1 Cloud Change



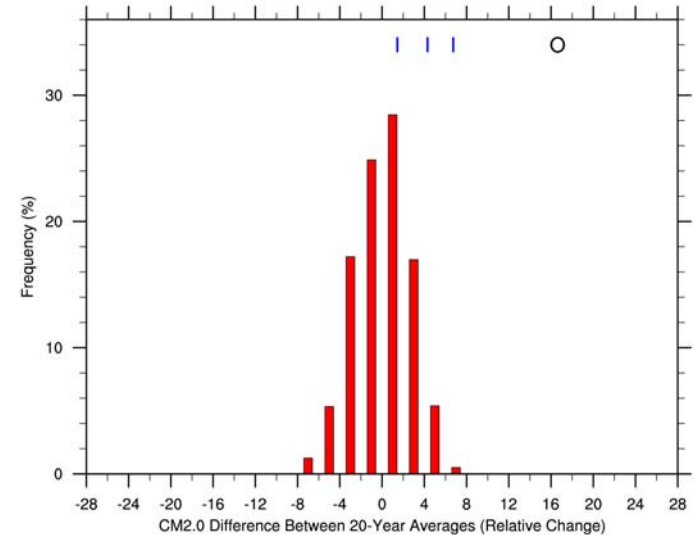
the frequency distribution of interdecadal cloud change in the coupled control runs will be examined in the specified region to determine if it is comparable to the observed upper-level cloud change

Central Eq. Pacific Control Comparison

CCSM3 Cloud Change



CM2.0 Cloud Change



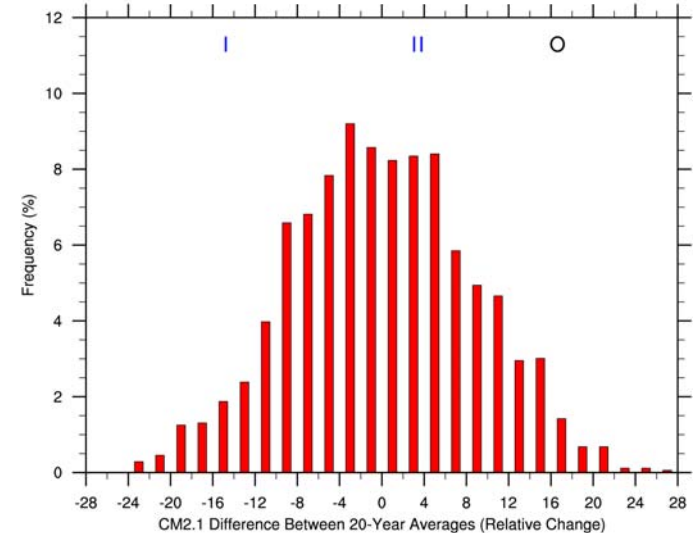
red = control run histogram

blue = IPCC 20th Century runs

green = AMIP runs

black = observed

CM2.1 Cloud Change



Conclusions

- Between 1957-1976 and 1977-1996...
observed upper-level cloud cover increased over the central equatorial Pacific
decreased over the adjacent subtropics
decreased over the western tropical Pacific
- Physically consistent changes occur in precipitation and surface atmospheric circulation.
- These changes are much larger than the expected linear response to the observed increase in Niño3.4 SST.

Conclusions

- CCSM3 AMIP simulations largely reproduce the observed central Pacific upper-level cloud changes, albeit with relatively weaker magnitude.
- CCSM3, CM2.0, and CM2.1 IPCC 20th Century runs do not reproduce the observed cloud changes.
- Interdecadal cloud variability in the CCSM3 and CM2.0 control runs is much weaker than the observed interdecadal cloud change.
- Interdecadal cloud variability in the CM2.1 control run has more realistic amplitude (*but interannual SST variability is too high*).

Conclusions

- The CCSM3, CM2.0, and CM2.1 20th Century simulations (if believable) suggest that the observed tropical cloud changes are not a climate response to anthropogenic forcing.
- Multidecadal tropical cloud variability is too weak in CCSM3 and CM2.0.